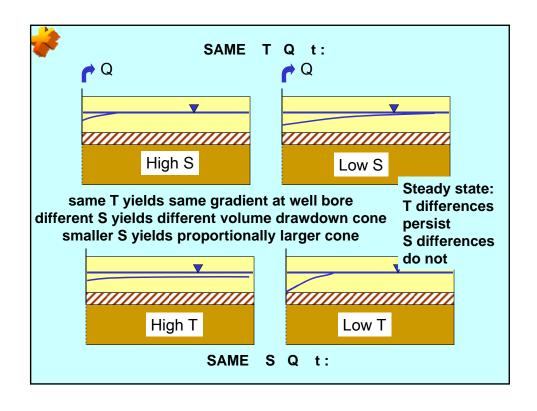
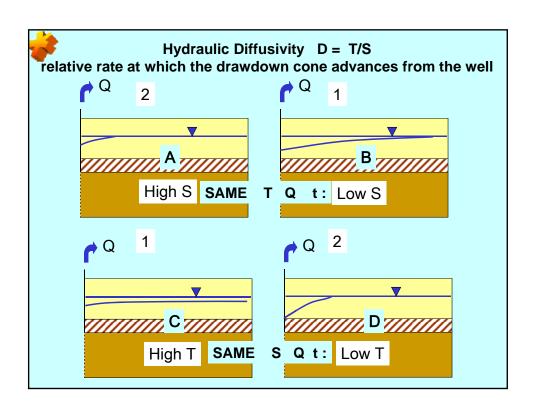
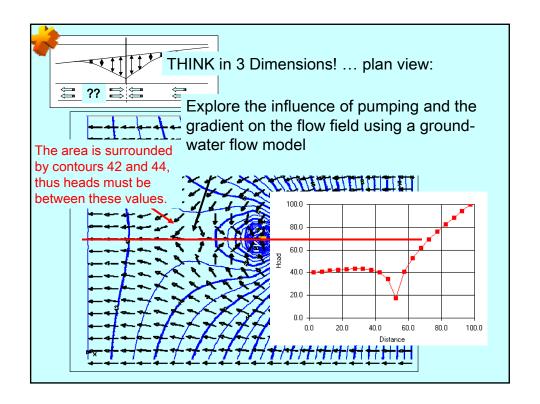
∳ Me	eters and	Davs					
Group 1	d	h _{max}	h _{x=12.5}	h _{x=37.5}	qriv	qcan	
K=0.01 w=0.000274	18.17	5.92	5.85	4.98	-0.00498	+0.00872	
Group 2 K=1x10 ⁻⁴ w=0.000274	24.93	41.58	36.13	36.00	-0.00683	+0.00687	
Group 3 K=1x10 ⁰ w=0.000274	-658.28	12.03	4.63	3.48	+0.18035	+0.19405	
Group 4 K=0.01 w= -0.000137	38.66	(min) 2.35	3.86	2.36	+0.00530	-0.00155	
Group 5 K=0.01 w=0.0	-infinity	infinity	4.62	3.46	+0.00187	+0.00187	
Group 5: q same as using Dupuit approx dh/dx with avg h for b							

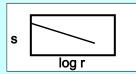
				Arrows			
100	Meters and	Days		Conceptual			
	d		can	Only .			
Group 1		•		^			
K=0.01	18.17	-0.00498	+0.00872 = 0.0137				
Recharge=	0.0137			← →			
Group 2							
K=1x10 ⁻⁴	24.93	-0.00683	+0.00687 = 0.0137	* \			
Recharge=0	0.0137			+ + \			
Group 3							
	-658.28	+0.18035	+0.19405 = 0.0137				
Recharge=0.0137							
Group 4				A			
K=0.01	38.66	+0.00530	-0.00155 = 0.0068	5			
Discharge=	0.00685			→ ←			
Group 5							
K=0.01	-infinity	+0.00187	+0.00187 = 0	+ +			
w=0.0	,						
Group 5: q same as using Dupuit approx dh/dx with avg h for b							
J. 53.P C			Land Shele and and and				







s vs. log r - is a straight line, if assumptions are met, drawdown decreases logarithmically with distance from the well because gradient decreases linearly with increasing area $(2\pi rh)$



$$Q = \frac{2\pi T (h_2 - h_1)}{\ln(r_2/r_1)}$$
 Theim Eqtn

 $T = transmissivity [L^2/T]$

Q = discharge from pumped well [L 3 /T]

r = radial distance from the well [L]

h = head at r [L]

and rearranging to get T from field data:

$$T = \frac{Q}{2\pi(h_2 - h_1)} \ln(\frac{r_2}{r_1})$$

Plot before applying equations

- to verify conditions are appropriate for application of equations
- to identify data problems

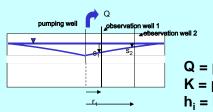
In an unconfined aquifer, T is not constant

If drawdown is small relative to saturated thickness, confined equilibrium formulas can be applied with only minor errors Otherwise call on Dupuit assumptions and use:

$$Q = \pi K \frac{(h_2^2 - h_1^2)}{\ln(r_2/r_1)}$$

or, to determine K from field measurements of head:

$$K = \frac{Q \ln(\frac{r_2}{r_1})}{\pi(h_2^2 - h_1^2)}$$



Q = pumping rate [L³/T] K = permeability [L/T]

h_i = head @ a distance r_i from well [L] using the aquifer base as datum

The aquifer base must be the datum because the head not only represents the gradient but also reflects the aquifer thickness, hence the flow area.